Evidence of dispersion in an artificial water-saturated sand sediment

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A laboratory experiment was conducted to measure the speed of sound in an artificial water-saturated granular sediment composed of cleaned and sorted medium-grained sand and degassed distilled water. The experiment was conducted within a range of frequencies where dispersion is predicted by a number of existing models. Between 2 and 4 kHz, the sound speed was inferred from measurements of the resonance frequencies of a thin-walled cylindrical container filled with the material. An elastic waveguide model was used to account for the effect of the finite impedance of the walls, although this effect was found to be small. From 20 to 300 kHz, the sound speed was obtained directly from time-of-flight measurements within the sediment. Dispersion in close agreement with the Williams effective density fluid model [K. L. Williams, J. Acoust. Soc. Am. 110, 2276-2281 (2001)] was observed. © 2007 Acoustical Society of America. [DOI: 10.1121/1.2404619]

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I. INTRODUCTION

A complete understanding of sound propagation in water-saturated granular sediments has not been achieved. A number of experimental studies have appeared in the literature that report either direct or indirect evidence of sound speed dispersion in this medium. Such results generally support the Biot-Stoll model (and its variants) for sound propagation. The Biot-derived models predict a low-frequency sound speed, followed by an increase to a higher sound speed beyond some transition frequency. There are also a number of experimental studies that found no evidence of sound speed dispersion and support Hamilton’s dispersionless fluid model. The dispersion predicted by the Biot models occurs over a range of about two orders of magnitude in frequency (from around 100 Hz to 10 kHz for medium-grained sand, for example). It is experimentally difficult to obtain measurements on a uniform body of sediment using a single technique that spans the entire range of interest. Further, the predicted dispersion is on the order of 10% and achieving experimental accuracy significantly greater than this has been difficult. Surface and volume inhomogeneities in natural sediments also cause uncertainty in experimental measurements.

A series of laboratory experiments designed to overcome some of these difficulties were performed on a water-saturated sand sediment. The sediment was contained within a thin-walled cylindrical vessel large enough to permit time-
of-flight sound speed measurements within the bulk sediment above 20 kHz. Lower frequency sound speeds were inferred from the measured resonance frequencies of the cylindrical sediment-filled container. For this analysis, transverse waves within the sediment were ignored. The sediment was modeled as a fluid, but the sediment sound speed was not constrained to be constant. The walls of the cylinder were initially approximated as a pressure release boundary (details of the assumption’s validity are discussed in the following). The cylinder was placed upon a layer of insulation foam, which yielded a pressure release lower boundary condition. The water/air interface provided a pressure release boundary at the top of the sediment. Therefore, a pressure release waveguide model was used to relate the measured resonance frequencies to the effective sediment sound speed. In order to identify individual modes, and hence infer frequency-dependent sound speeds from resonances, both symmetric and asymmetric excitation of the sediment cylinder was used. The spatially dependent phase of the acoustic pressure, relative to the drive signal, was employed to aid in mode identification. Dispersion was observed and these results are presented and compared to the effective density fluid model, hereafter referred to as EDFM. The effect of the finite impedance imparted by the cylindrical wall of the sediment container is also discussed. A model of an elastic-walled waveguide was used to relate the sediment sound speeds that were inferred from axisymmetric modes in the tank to the values that would be observed in an unconfined free field. The effect was found to be small (less than 0.53%) at the lowest experimental frequency and decreased as the frequency increased.

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A laboratory experiment was conducted to measure the speed of sound in an artificial water-saturated granular sediment composed of cleaned and sorted medium-grained sand and degassed distilled water. The experiment was conducted within a range of frequencies where dispersion is predicted by a number of existing models. Between 2 and 4 kHz, the sound speed was inferred from measurements of the resonance frequencies of a thin-walled cylindrical container filled with the material. An elastic waveguide model was used to account for the effect of the finite impedance of the walls, although this effect was found to be small. From 20 to 300 kHz, the sound speed was obtained directly from time-of-flight measurements within the sediment. Dispersion in close agreement with the Williams effective density fluid model [K. L. Williams, J. Acoust. Soc. Am. 121 (2), February 2007].
II. DESCRIPTION OF THE EXPERIMENT

A. Sediment container

The sediment was contained in a right-circular-cylindrical tank possessing walls made of high-density polyethylene (HDPE). The mean inner diameter was 0.686 m, the height of the sediment within the vessel was 0.80 m, and the mean wall thickness was 0.48 cm. The tank, which had a flat bottom, was placed upon a 10-cm-thick styrofoam panel which, in turn, rested on a concrete floor.

The sediment was prepared in the following manner. Filtered fresh water was added to the tank and degassed in situ by heating it to 60 °C using several immersion heaters. Circulation was imposed during this process with a motorized stirring impeller. The water was kept at 60 °C for several hours during which time the dissolved gas concentration attained equilibrium. Then, approximately 500 kg of a commercial silica blasting sand was slowly added to the water, as the stirring continued. The sediment was then allowed to cool to room temperature, over a period of 24 h. As the system cooled, the water phase became undersaturated with gas and thus drew into solution any gas pockets entrained by the introduction of the sand, resulting in a fully saturated sediment. A sieve analysis conducted on approximately 500 ml of the sand yielded grain sizes ranging from 150 to 600 μm, with a mean size of 300 μm.

B. High-frequency time-of-flight measurements

The sound speed within the sediment was measured using a time-of-flight technique at 20, 50, 100, and 300 kHz. A schematic diagram of this experiment is shown in Fig. 1. The signal source was a function generator programmed to produce sinusoidal pulses of various length, depending upon the frequency, to avoid multipath interference. The source signal was bandpass filtered (24 dB/octave) at the pulse center frequency with a Krohn-Hite model 34A filter and then amplified by a Krohn-Hite model 7500 wide-band power amplifier and directed to a Bruel & Kjær model 8103 miniature hydrophone, operating as a source (labeled S in Fig. 1). Two more 8103 hydrophones were positioned 10.0 and 17.6 cm away from the source on a horizontal line (parallel to the sediment surface), at a depth of 15.0 cm. The latter two hydrophones acted as receivers (labeled R1 and R2 in Fig. 1). The relative positions of the source and receivers were constrained to within ±0.25 mm by mounting each hydrophone and its cable inside a stiff stainless steel tube. These tubes were mounted to a steel frame and fastened securely in place. The R1–R2 separation distance was set and measured with a micrometer caliper prior to inserting the hydrophones into the sediment. The rigidity of the mounting scheme was sufficient to prevent caliper-measurable deflection during the inserting process. The hydrophone-enclosing tubes were filled with degassed water to minimize acoustic disturbance of the sediment.

The signals from R1 and R2 were conditioned with a Bruel & Kjær model 2692 charge amplifier, then bandpass filtered (24 dB/octave) at the pulse center frequency with Krohn-Hite model 34A filters. The filtered source and receive signals were acquired with a digital oscilloscope. At each frequency, 100 pings were collected and the averaged signals were sent to a computer for processing. The R1 and R2 signals were cross-correlated to obtain the time-of-flight and then sound speed was calculated from the known sensor separation distance. The uncertainty in sound speed was calculated from the previously stated position uncertainty, and the digitization sample rate and found to be between ±6.6 and ±9.9 m/s at 300 and 20 kHz, respectively.

C. Low-frequency resonance measurements

The apparatus for this portion of the experiment is shown in Fig. 2. A custom-fabricated Tonpilz-type transducer was used as a source. It is described in Ref. 17 and has a circular-shaped radiating face with an accelerometer imbedded in it, and therefore can be driven at constant velocity amplitude by suitable monitoring of the accelerometer signal and modification of the excitation voltage amplitude. The source was placed face-down on the surface of the sediment in one of two positions, to be described later. The source signal was a band-limited (2–5 kHz) periodic chirp, provided by a Hewlett-Packard 89410A vector signal analyzer (VSA) and amplified by a Crown CE 1000 power amplifier. The chirp amplitude was modified onboard the VSA to produce a source velocity amplitude that was uniform across the frequency range, as discussed earlier. The acoustic pressure was measured in the sediment with B&K 8103 hydrophones at two positions, labeled R1 and R2 in Fig. 2. The depth was 6.7 cm and each sensor was positioned 10 cm from the vessel wall. When observed from above, R1–S–R2 formed a straight line. The signals from R1 and R2 were conditioned.
by the B&K 2692 charge amp, which included bandpass filtering from 10 to 10 kHz, and digitized by the VSA. The complex spectrum of each hydrophone signal (with phase referenced to the excitation signal) was calculated onboard the analyzer using 100 spectral averages and converted to complex acoustic pressure spectra using the calibrated hydrophone sensitivities.

III. ACOUSTIC MODEL OF SEDIMENT AND CONTAINER

An acoustic model of the sediment/container system is needed to infer sound speeds from measured resonance frequencies. In this analysis, the sediment was considered to be an effective fluid with wave number \( k \), which is a valid approach for granular sediments.\(^1\) We assume the phase speed can be frequency dependent, but \( \text{Im}[k] \ll \text{Re}[k] \), so that the resonance frequencies depend primarily upon the phase speed and have weak dependence upon the attenuation. Transverse waves in the sediment were ignored.

Initially, we ignore the container walls and consider the sides of the cylinder to have a pressure-release boundary condition under the following justification. (1) Typical values of the specific acoustic impedance of the materials involved are given in Table I. Based on these values and the wall thickness (0.48 cm), the three-medium reflection coefficient \( R \), given by Eq. 6.3.7 in Ref. 18 for pressure waves incident upon the wall, is \( R = 1.00 \pm 1.5^\circ \) between 2 and 5 kHz. (2) This was qualitatively verified by the observation of inverted reflections from the walls during the time-of-flight measurements, and (3) by the absence of plane-wave longitudinal mode resonances in the system. The air-sediment interface at the top of the container is clearly pressure release. Finally, the container was placed upon a 10-cm-thick styrofoam building insulation panel to approximate a pressure release boundary condition on the bottom. For the cylindrical coordinate system shown in Fig. 1 and sinusoidal excitation, the acoustic pressure in a right-circular cylinder with \( p=0 \) boundary conditions at all surfaces is

\[
P_{mnN}(r, \theta, z, t) = A_{mnN} \alpha_{mn} \tau^\text{out} \times \cos(m \theta - \theta_{0m}) \sin(N \pi z/h) e^{i\omega t},
\]

and has resonance frequencies

\[
f_{mnN} = \frac{c}{2} \sqrt{\left( \frac{\alpha_{mn}}{\pi a} \right)^2 + \left( \frac{N}{h} \right)^2},
\]

where \( \alpha_{mn} \) is the \( n \)th root of the Bessel function \( J_m(x) \), \( c \) is the effective speed of sound in the sediment, \( a \) is the tank radius, \( h \) is the sediment height, \( A_{mnN} \) is the modal amplitude coefficient, and \( \theta_{0m} \) is a phase angle which depends on the source condition. Equations (1) and (2) were adapted from Ref. 19. A plot of resonance frequencies versus longitudinal mode number \( N \) is shown in Fig. 3 for the six lowest-order mode families. Five modes, in order of ascending frequency, are labeled with roman numerals from I to V. Note that above mode III, multiple modes appear with similar eigenfrequencies.

Calculation of sound speed from a measured spectral peak requires one to identify the specific eigenmode and its \( m, n, N \) values. Mode family \( 0,1, N \) is axisymmetric and mode family \( 1,1, N \) is asymmetric about the center of the sediment cylinder, as shown in Fig. 4. These two modes were selectively excited by using the source positions shown in Fig. 4. To excite mode family \( 0,1, N \), the source was placed at the center of the sediment cylinder. To excite mode family \( 1,1, N \), the source was placed halfway between the center and the container wall. The relative phase of the acoustic pressure observed at R1 and R2 was used to help identify the mode associated with an observed resonance peak. For example, Eq. (1) yields acoustic pressures at R1 and R2 that are in phase for modes I, II, and V, and 180° out of phase for modes III and IV.

![FIG. 3](image-url) The resonance frequencies for the six lowest-order mode families for a constant sound speed of 1730 m/s. The families are labeled \( m,n \) around the top and right-hand side of the plot. The plane wave mode would be \( 0,0, N \) if it existed.

![FIG. 4](image-url) The mode shapes given by Eq. (1) for two mode families. The source positions for each case are labeled S. The receiver positions, which remain the same in both cases, are labeled R1 and R2.

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IV. RESULTS

A total of four averaged spectra were obtained during the resonance frequency experiment. These consisted of the acoustic pressure measured at R1 and R2, for both of the source positions shown in Fig. 4. The pressure spectra obtained at R1 are shown in Fig. 5. For the spectral peaks labeled I, II, and V, there was less than one degree phase difference between R1 and R2. These three spectral peaks are associated with mode family 0,1,1 and correspond with the symmetric eigenmodes labeled I, II, and V in Fig. 3. Returning to Fig. 5, the spectral peaks labeled III and IV exhibited an R1–R2 phase difference of 170°. These peaks correspond to the asymmetric modes III and IV, shown in Fig. 3 and associated with mode family 1,1,1. Effective sediment sound speeds \( c \) were calculated from each of the identified peaks in Fig. 5 using Eq. (2). Although not shown, similar results were obtained from the sensor at R2.

Both the resonance-based and time-of-flight-based sound speeds are shown in Fig. 6. The speeds are normalized by the speed of sound in distilled water at the ambient temperature (23.5±0.5 °C). The sound speed predicted by the effective density fluid model is also shown in Fig. 6. The pertinent equations of the model are shown in the following for convenience. The EDFM considers a water-saturated granular sediment with a total mass density \( \rho \), pore fluid mass density \( \rho_f \), sediment particle mass density \( \rho_s \), individual sediment grain bulk modulus \( K_r \), pore fluid bulk modulus \( K_f \), porosity \( \beta \), tortuosity \( \alpha \), permeability \( \kappa \), pore fluid viscosity \( \eta \), and complex sound speed \( c \). The sound speed is a function of angular frequency \( \omega \) and is given by

\[
c = \sqrt{\frac{\omega \rho_f}{c_{\text{eff}}(\omega)}},
\]

where the compressibility is determined by a mixture rule

\[
H = \left( \frac{1 - \beta \rho_s}{K_f} + \frac{\beta}{K_f} \right)^{-1}.
\]

and the effective density is

[Table II. Material parameters used in the effective density fluid model (SI units). Parameters obtained from measurements on the sediment are in BOLDFACE. The mean value is followed by the range of variation in multiple measurements. The remaining parameters are from Ref. 12.]

<table>
<thead>
<tr>
<th>Property</th>
<th>Value</th>
</tr>
</thead>
<tbody>
<tr>
<td>Porosity</td>
<td>0.370 (0.368–0.371)</td>
</tr>
<tr>
<td>Sand grain mass density</td>
<td>2650</td>
</tr>
<tr>
<td>Water mass density</td>
<td>998</td>
</tr>
<tr>
<td>Sand grain bulk modulus</td>
<td>3.6 X 10^10</td>
</tr>
<tr>
<td>Water sound speed</td>
<td>1492</td>
</tr>
<tr>
<td>Water viscosity</td>
<td>0.001</td>
</tr>
<tr>
<td>Permeability</td>
<td>4.58 X 10^-11 (4.18 X 10^-11 – 4.74 X 10^-11)</td>
</tr>
<tr>
<td>Tortuosity</td>
<td>1.25</td>
</tr>
</tbody>
</table>

V. SEDIMENT SOUND SPEED FROM THE EFFECTIVE DENSITY FLUID MODEL

The effective density fluid model of sound propagation in water-saturated granular sediments is described in detail elsewhere, but the equations for prediction of the sediment sound speed are repeated here for convenience. The EDFM considers a water-saturated granular sediment with a total mass density \( \rho \), pore fluid mass density \( \rho_f \), sediment particle mass density \( \rho_s \), individual sediment grain bulk modulus \( K_r \), pore fluid bulk modulus \( K_f \), porosity \( \beta \), tortuosity \( \alpha \), permeability \( \kappa \), pore fluid viscosity \( \eta \), and complex sound speed \( c \).
TABLE III. Material parameters used in the evaluation of Eq. (A1) (SI units).

<table>
<thead>
<tr>
<th>Material Parameter</th>
<th>Value</th>
</tr>
</thead>
<tbody>
<tr>
<td>Sediment mass density, $\rho_s$</td>
<td>2039</td>
</tr>
<tr>
<td>HDPE mass density, $\rho_{HDPE}$</td>
<td>1000</td>
</tr>
<tr>
<td>HDPE compressive sound speed, $c_t$</td>
<td>2500</td>
</tr>
<tr>
<td>HDPE transverse sound speed, $c_t$</td>
<td>680</td>
</tr>
</tbody>
</table>

The parameter $F$ represents a deviation from Poiseuille flow of the fluid within the pores and is given by

$$
F(\varepsilon) = \frac{\rho_f}{\rho_f + \rho_s} \left( \frac{\alpha(1 - \beta)\rho_s + \beta(\alpha - 1)\rho_f + i\beta_F \eta}{\rho_w \omega_k} \right) \left( \frac{\beta(1 - \beta)\rho_s + (\alpha - 2\beta + \beta^2)\rho_f + i\beta_F \eta}{\omega_k} \right).
$$

(5)

The relative size of the measurement uncertainty due to tank dimension uncertainty is compared to the finite wall impedance correction in Fig. 7. For clarity, this comparison is only shown for one of the data sets. The circle data points from Fig. 6 are repeated in Fig. 7 and their values after correction for the finite wall impedance are shown. The finite wall impedance biases the axisymmetric resonance measurements (circles and squares in Fig. 6) by $+0.57\%$ or less, but this is significantly less than the measurement uncertainty due to the tank dimensions.

An exact analytic model for asymmetric excitation of a liquid-filled waveguide with walls of arbitrary thickness is not available. We used an approximate model$^{23}$ which treats the wall as a shell and yields a dispersion relation Eq. (A6) for both axisymmetric and asymmetric modes. Otherwise, the corrections procedure is the same as for the axisymmetric case, discussed earlier. The dimensions of the sediment tank

FIG. 7. The effect of ignoring the finite impedance of the tank walls is shown for the axisymmetric modes. The large circles are the sediment sound speeds inferred from a pressure-release model of the system, repeated here from Fig. 6, and the error bars represent the measurement uncertainty due to imperfect knowledge of the tank dimensions. The small circles are the sediment sound speeds inferred from a model which accounts for the finite impedance of the walls of the tank. Ignoring the finite impedance of the walls results in a maximum bias error of $+0.53\%$ at the lowest frequency measured.
The solid curve is the EDFM prediction based on the mean porosity and except that the resonance method measurements were corrected for the finite wall impedance of the tank. The extent of measurement uncertainty for the range of the porosity measurements and the standard deviation of the permeability measurements. The dashed curves above and below the solid curve represent the maximum and minimum EDFM predictions associated with the range of the porosity measurements and the standard deviation of the permeability measurements.

are well within the limits required by the shell approximation. To ensure continuity, we compared the phase speeds calculated by Eqs. (A1) and (A6) for the symmetric mode 0,1 and found excellent agreement. We are therefore confident that the approximate Eq. (A6) is accurate for the geometric and material properties used here. The data and model prediction presented in Fig. 6 are repeated in Fig. 8, but with the finite wall impedance correction applied to all the data from both axisymmetric and asymmetric modes. The asymmetric data before and after correction are presented in Table V, and shows a maximum correction of +0.76%. Note that the time-of-flight measurements are not biased by the presence of the tank walls because the environment appears to be a free field for the pulse lengths and frequencies used.

VII. CONCLUSIONS

Laboratory sound speed measurements were obtained in a water-saturated granular sediment composed of commercial silica blasting sand and fresh water. Above 20 kHz, a direct time-of-flight method was used and between 2 and 5 kHz, an indirect method was used to infer frequency-dependent sound speeds from measured resonance frequen-

cies. The pressure-release approximation yields a surprisingly good model for a thin-walled plastic tank filled with a water saturated sediment, as shown in Figs. 6–8. There is at maximum about 0.5% error in sound speeds inferred from the pressure release model.

Considering the more realistic analysis which does account for the finite impedance of the tank walls and is shown in Fig. 8, the EDFM model\(^2\) describes the sound speed measurements very well. Most of the low-frequency measurements obtained with the resonance method are just lower than the predicted values, but the agreement between model and measurement is well within the measurement uncertainty. Three of the four high-frequency time-of-flight measurements are just higher than the prediction, but again, within the range of measurement uncertainty. The overall conclusion is that dispersion is clearly observed, as predicted by the Biot-derived models, and that the EDFM accurately describes the sound speed observed in this artificial laboratory sediment. Note that within the experimental frequency range, both measurements and model indicate about 4% dispersion. Evaluation of the EDFM from its low frequency asymptote (\(\approx 100\) Hz) to its high frequency asymptote (\(\approx 1\) MHz) yields a total of about 8% dispersion. A secondary conclusion can be drawn from the spread in the resonance-inferred data, which sampled the entire sediment volume, as compared to the high frequency measurements which only sampled a small portion of the volume. Even in a carefully prepared laboratory specimen, granular sediments are acoustically nonhomogeneous.

Overprediction of sediment sound speed by Biot-type models in the transition region between the low and high frequency asymptotes has been previously reported in Refs. 1 and 4. One explanation for this deviation in the previous work is the existence of undissolved air. In the present experiment, significant effort was expended to achieve a gas-free sediment, and it appears that if there was any remaining undissolved air (in microscopic crevices for instance) the effect was not acoustically significant.

ACKNOWLEDGMENTS

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APPENDIX

1. Axisymmetric dispersion relation

We considered sound propagation along the longitudinal axis of a lossless fluid cylinder enclosed by a lossless, arbitrary thickness, elastic cylindrical wall. Acoustic energy at angular frequency \(\omega = 2\pi f\) propagating in this system (where \(f\) is the number of cycles per second) is governed by a set of exact wave theoretic equations given originally by Del Grosso\(^21\) and refined by Lafleur and Shields.\(^22\) The dispersion relation for this system is given in Eq. (A1). The ele-
ments of Eq. (A1) can be complex valued, but real values of \( q_{0m} = \omega / c_{0m} \) that cause the left-hand side of Eq. (A1) to become zero lead to propagating modes that have phase speed \( c_{0m} \). At any given frequency there can be multiple permissible modes, which are indexed by integer values of \( m \). Since the phase speed is frequency-dependent, at a given sediment cylinder resonance frequency \( f_{0m,N} \), the phase speed will attain a value denoted by \( c_{0m,N} \). The zero subscript that precedes the index \( m \) indicates that these are axisymmetric modes. The dispersion relation is

\[
1 + [L_{11}(P_m)L_{00}(T_m)] \left( \frac{\pi^2 q_{0m}^2 b d P^2 q^2}{8 E_m} \right) + [L_{11}(T_m)L_{00}(P_m)] \left( \frac{\pi^2 b d E_m}{8 q_{0m}^2} \right) + [L_{10}(P_m)L_{01}(T_m) + L_{01}(P_m)L_{10}(T_m)] \\
\times \left( \frac{\pi^2 b d P^2 q^2}{8} \right) + [b L_{11}(P_m)_{L_{10}(T_m)} + d(1 + Q_m b) L_{11}(P_m)_{L_{01}(T_m)}] \left( \frac{\pi^2 P_m^2 T_m}{8 E_m} - \frac{\pi^2 P_m^2 q_{0m}^2 T_m}{8 E_m} \right) + [b L_{11}(T_m)L_{10}(P_m) + d(1 + Q_m b) L_{11}(T_m)L_{01}(P_m)] \left( \frac{\pi^2 P_m}{8 q_{0m}^2} - \frac{\pi^2 P_m^2 q_{0m}^2}{8 E_m} - \frac{\pi^2 P_m^2}{4 E_m} \right) = 0,
\]

(A1)

where the following definitions have been used:

\[
X_{0m} = b \sqrt{k_i^2 - q_{0m}^2}, \quad P_m = \sqrt{k_i^2 - q_{0m}^2}, \quad T_m = \sqrt{k_i^2 - q_{0m}^2},
\]

(A2)

\[
q_{0m} = \omega / c_{0m}, \quad k_i = \omega / c_j, \quad k_i = \omega / c_i, \quad \Omega = \omega / c_i,
\]

(A3)

\[
E_m = \frac{q_{0m}^2 - k_i^2/2}{2}, \quad Q_m = \frac{\rho_l \omega^2 b J_i(X_{0m})}{2 \mu_s c^2 X_{0m} J_i(X_{0m})},
\]

(A4)

\[
L_{mn}(\xi) = J_n(\xi) Y_n(b \xi) - J_n(b \xi) Y_n(\xi),
\]

(A5)

and the geometric and material input parameters are: \( c_1 \) is the intrinsic sound speed in the inner liquid (the sound speed that the sediment would exhibit in the free field), \( c_j \) and \( c_i \) are the longitudinal and transverse (shear) speed for the wall material, \( \rho_l \) and \( \rho_w \) are the liquid and wall material densities, and \( b \) and \( d \) are the inner and outer radii of the pipe, respectively. \( J_n(x) \) and \( Y_n(x) \) are Bessel functions of the first and second kind, respectively.

2. Asymmetric dispersion relation

We modeled nonaxisymmetric sound propagation along the longitudinal axis of a lossless fluid cylinder enclosed by a lossless, finite thickness, shell-like cylindrical wall, using the formulation of Fuller and Fahy. The dispersion relation is given in terms of a determinant

\[
\det[L_{ij}] = 0 \quad \text{(A6)}
\]

where

\[
L_{11} = -\Omega^2 + (k_{mn} b)^2 + \frac{1}{2} (1 - \nu) n^2, \quad L_{12} = \frac{1}{2} (1 + \nu) n (k_{mn} b), \quad L_{13} = \nu (k_{mn} b), \quad L_{21} = L_{12}, \quad L_{22} = -\Omega^2 + \frac{1}{2} (1 - \nu) (k_{mn} b)^2 + n^2, \quad L_{23} = n,
\]

(A7a)

In Eq. (A7), \( \Omega = \omega / c_i, \nu \) is the ratio of Poisson of the wall material, \( \beta^2 = h^2 / 12 b^2 \), \( h \) is the wall thickness, and FL represents fluid loading

\[
\text{FL} = \Omega^2 (p_{w} / \rho_{w})(b / h) (1 / k_{mn} b) (J_n(k_{mn} b) / J_n'(k_{mn} b)),
\]

(A8)

The radial fluid wave number \( k_{mn} b \) and the axial wave number \( k_{mn} b \) are related by

\[
k_{mn} b = [\Omega^2 (c_1 / c_i)^2 - (k_{mn} b)^2]^{1/2},
\]

(A9)

3. Illustration of finite wall impedance correction scheme

The effect of the finite wall impedance and the method of correction for this effect are demonstrated in Fig. 9 for the axisymmetric mode. The phase speed for the 0,1,1 mode is shown for both the pressure release model and the elastic waveguide model Eq. (A1). Note that the predicted phase speeds are similar, therefore we do not expect the corrections to be large. For a given experimentally observed resonance frequency \( f_{0,1,1} \), the observed modal phase speed is calculated from

\[
c_{0,1,1} = \lambda f_{0,1,1},
\]

(A10)

where \( \lambda = 2 h / N \). Then the effective sediment sound speed \( c_1 \) is varied in Eq. (A1) until the predicted modal phase speed is \( c_{0,1,1} \). This procedure was repeated for each resonance frequency identified from an axisymmetric mode. The correction procedure is identical for the symmetric modes, except Eq. (A1) is replaced by Eq. (A6). The results are reported in Table IV and shown in Fig. 8.
The results of the correction are shown in Table IV.

The vertical axis is normalized by the sediment effective intrinsic sound speed $c_1$. The basis for the procedure used to correct for the finite wall impedance is illustrated for the one of the data sets (the open circles of Fig. 6). Integral multiples of one-half wavelength occupy the sediment cylinder along its height for the 0, 1, 2, 3, 4. Effect of neglecting viscosity in the wall effective sediment sound speed $c$, that would be measured in the free field. Knowing of these mode shapes from Eq. (1), the measured resonance frequencies $f_{0,1,2}$ and Eq. (A10) allow one to use Eq. (A1) to relate the measured value of $c_{0,1,2}$ to the intrinsic effective sediment sound speed $c_1$ that would be measured in the free field. The results of the correction are shown in Table IV.

4. Effect of neglecting viscosity in the wall impedance correction

The above-presented elastic waveguide models approximate the internal fluid as inviscid, therefore the fluid can slip at the boundary with the wall. In this section it is shown that neglecting the sediment viscosity in the waveguide models has an insignificant effect on the phase speeds in this experiment. This is due to the large size of the sediment tank in comparison to the boundary layer at the wall.

In this work, the sediment is approximated as an effective fluid and can be assigned an effective viscosity with respect to pipe wall interaction, as is done in the modeling of slurry flows, and in acoustic studies of sediments in tubes. For mixtures of liquids and sand with high solid volume concentrations (such as the present sediment) the effective mixture viscosity $\mu_{\text{mix}}$ is dependent upon the solid volume concentration $\theta$ as

$$\frac{\mu_{\text{mix}}}{\mu} = 1 + 2.5 \phi + 10.05 \phi^2 + 2.73 \times 10^{-3} \exp(16.6 \phi),$$

where $\mu$ is the viscosity of the host liquid. The sediment porosity in this work is 0.37 therefore $\phi=0.63$ and Eq. (A11) yields a viscosity ratio $\mu_{\text{mix}}/\mu=102$.

For acoustic waves propagating within a viscous fluid inside a pipe, an acoustic boundary layer is formed at the wall. The resulting acoustic effects can be modeled with a complex wave number

$$k = \frac{\omega}{c} + (1-i)\alpha_{\text{wall}},$$

where $c$ is the effective sound speed of the sediment inside the waveguide and $\omega/c \gg \alpha_{\text{wall}}$ is assumed. Thermal losses are considered negligible compared to the viscous losses, which are characterized by

$$\alpha_{\text{wall}} = \frac{1}{a} \sqrt{\frac{\mu_{\text{mix}} \omega}{2pc^2}},$$

where $a$ is the inner radius of the pipe and $p$ is the total mass density of the sediment. Equation (A12) indicates a reduced phase speed inside the waveguide due to the viscous boundary layer. The effect of the viscosity on phase speed can be expressed in a ratio

$$\frac{\bar{c}}{c} = \left(1 + \frac{\alpha_{\text{wall}}}{\omega}ight)^{-1},$$

where $\bar{c} = \text{Re} [\omega/k]$ is normalized by the sound speed in absence of a viscous boundary layer, $c$.

Using the sediment mass density from Table III, the highest frequency in the resonance regime of the experiment (3788 Hz) and the mean sound speed observed inside the waveguide at the same frequency (1740 m/s), Eq. (A14) yields $\bar{c}/c=0.99905$, which is a negligible systematic error in comparison to other sources of measurement uncertainty in this experiment and is ignored.