

Fluid Mud in Energetic Systems: FLUMES II

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LONG-TERM GOALS

The goal of this research is to develop greater understanding of the dynamics of fluid mud and its role in the transport and deposition of sediment on the continental margin. In particular, we seek greater understanding of the processes that influence the formation and maintenance of fluid mud in energetic environments.

OBJECTIVES

The research is a process-based study that addresses three primary objectives:

- Determine controlling factors in the formation and destruction of fluid mud under a sheared flow
- Verify Richardson number dependence for suppression of turbulence and carrying capacity of a high-concentration suspension
- Evaluate effects of mixed grain size on high-concentration suspensions

APPROACH

To evaluate the controls on fluid-mud formation and the influence of sediment-induced stratification on flow, in situ measurements of vertical gradients of velocity, suspended-sediment concentration, and fluid density throughout the water column, including the fluid-mud layer are necessary. In addition to these parameters, the thickness of the fluid mud layer through accelerating and decelerating flows must be measured.

The Peticodiac River located in the Upper Bay of Fundy was selected as the study site. The Peticodiac Estuary is a macro-tidal environment that has been modified by the construction of a causeway in the

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late 1960s. During construction, depositional rates downstream of the causeway were on the order of 1 cm day^{-1} resulting in a decrease in the cross-sectional area of up to 90% within one year of closure. Maximum tidal currents are now on the order 2 m s^{-1} and suspended-sediment concentrations regularly exceed 10 g l^{-1} and can reach 300 g l^{-1} (Curran et al. 2004). At slack tide, sediment settles rapidly forming fluid mud layers on the order of 1-2 m thick.

The Peticodiac River Estuary now provides an ideal natural lab where the formation and destruction of fluid mud can be studied under a sheared flow with a quasi-steady current and without the complicating effect of waves. Flow speeds can be varied by timing the experiments to coincide with the spring neap tidal cycle and by modifying the release of freshwater through the control structure. Measurements are carried out from a bridge that provides a solid framework for instrument deployment thus eliminating many of the challenges of shipboard operations in high-current environments. The Peticodiac has a distinct advantage over laboratory studies in terms of scale and avoids the practical issues of dealing with the quantities of mud required to create a fluid-mud layer and maintaining flume equipment at these high concentrations of suspended sediment.

For water-column sampling in high flow conditions a streamlined underwater buoyancy (SUBS) package designed at the Bedford Institute of Oceanography (BIO) was modified for use as an instrument profiler (Fig. 1). Current shear is determined using two Marsh McBirney electromagnetic current meters mounted 60 cm apart on the front of the package. An Ocean Sensors CTD and an optical backscatterance sensor (OBS) are mounted through the front and side of the package respectively. A sample port is co-located with the OBS and connected to a series of pumps to provide suspended sediment for OBS calibration and grain size analysis. To determine the depth of the lutocline and continuously monitor flow velocities, a dual frequency Knudsen echo sounder and an RD Instruments ADCP are mounted on a surfboard tethered to the bridge next to the location of the profiling package. Instruments are deployed from specially designed davits mounted on the bridge deck. Measurements are made every 45-60 minutes over tidal cycles.

All work is being conducted collaboratively between Tim Milligan of the Bedford Institute of Oceanography (BIO) and Gail Kineke of Boston College (BC). Brent Law (BIO) and graduate students Kristy Heath and Michelle Lermon (BC) provide technical support in the laboratory and field.

WORK COMPLETED

FLUMES I: Design and construction of the initial instrument package and davits was completed in May 2006. Two field experiments were carried out over two 1-week periods in June and August 2006. The June experiment was used to test equipment and observe high flow conditions. The August experiment coincided with a spring tide during a period of maximum seasonal sediment deposition in the Peticodiac.

Data analysis of the 2006 data is complete and results have been presented at the AGU Fall meeting, December 2006, and at the INTERCOH conference in Brest France, September 2007. Kristy Heath's Masters thesis is undergoing final revisions and her work was presented at the Estuarine Research Federation (now Coastal and Estuarine Research Federation) meeting in November 2007. Publications are in progress. Complete interpretation requires observation of the breakdown of the lutocline scheduled for October 2008.

FLUMES II: The work in 2008 has focused on preparation for the next field experiment scheduled for October 2008. The objective for this experiment is to capture intermediate flow conditions that should produce sufficient flow on the ebb to cause the resuspension of the fluid mud layer. The disappearance of Ocean Sensors, Inc. resulted in the inability to acquire spare parts for the Ocean Sensors CTD which served as the main controller and logger for the SUBS. Thus, a significant redesign of the profiling package was required. The Ocean Sensors CTD was replaced with an RBR CTD and a new pump controller was designed and built at BIO. A second SUBS was modified to hold an ADV and OBS5.

In preparation of the upcoming field experiment, a laboratory pre-calibration was performed comparing the response of different optical sensors for redundancy in the field. The sensors included two D&A Instruments OBS5s (internally recording high concentration sensors), a D&A Instruments OBS3 logged with an RBR CTD, two Seapoint Turbidity Sensors logged with RBR CTDs. Field calibration of an OBS3 using the *in situ* sediment samples collected was completed in 2007 yielded response up to $\sim 250 \text{ g l}^{-1}$. This calibration takes advantage of the non linear response of the OBS at concentrations greater than $\sim 5 \text{ g l}^{-1}$ that was demonstrated for fluid mud on the Amazon Shelf (Kineke and Sternberg 1995). The OBS5s (internally recording with two detectors) respond from 0-40 g l^{-1} Seapoint sensors to $\sim 20 \text{ g l}^{-1}$.

Another goal for the October experiment is to make direct turbulence measurements over a range of suspended sediment concentrations. ADVs are routinely used in fine-grained sedimentation and transport studies, but the performance in very high sediment concentrations (1-100+ g l^{-1} , i.e. fluid mud) has not been well documented. D. Velasco of Sontek/YSI has used samples provided by Kineke to test the response of ADVs at fluid mud concentrations. His results are that for low velocity ranges (long pulse lags), valid velocity data can be collected up to $\sim 50 \text{ g/l}$, $\sim 28 \text{ g/l}$, and $\sim 18 \text{ g/l}$ for the 5, 10 and 16 MHz ADVs, respectively (Velasco and Huhta 2007). We will further test the ADV response under more energetic conditions.

RESULTS

The June 2006 experiment occurred during a period of unusually high outflow of fresh water due to a period of prolonged rainfall prior to sampling. The August 2006 experiment occurred during a period of low outflow. Thus measurements were made during two end-member conditions: high outflow of fresh water with moderate concentrations (generally less than 10 g l^{-1}) and tidal flows of saltwater, high concentration suspensions (often greater than 10 g l^{-1} and up to 366 g l^{-1}).

An example from the August 2006 experiment shows water column profiles of density (due to salinity and temperature) and suspended-sediment concentration from early flood to early ebb with lower fluid density at the bottom that is compensated by the high suspended-sediment concentrations of the suspension (Fig.2). Approximately 1.5 hours after the passage of the tidal bore (at $\sim 15:00$) and a fully mixed turbulent flow, the water column begins to stratify and a high-concentration bottom layer forms persisting through the ensuing ebb (Fig.3). Measured suspended-sediment concentrations reached 286 g l^{-1} and low shear rates of 0.13 s^{-1} in the upper water column increased to $\sim 0.50 \text{ s}^{-1}$ through the lutocline 1 m above the bed, and decreased to ~ 0 within the fluid mud. At concentrations above 10 g l^{-1} the gradient Richardson number was consistently greater than $\frac{1}{4}$ suggesting suppression of turbulence (Fig. 4).

The threshold condition of free-stream velocity to the integrated buoyancy anomaly ($u_\infty/B^{1/2} < 2$; Trowbridge and Kineke 1994) for the formation of a high-concentration bottom layer was tested for all

profiles during the August experiment (Fig. 5). In general, the threshold condition is met for the profiles during decelerating conditions coinciding with the formation of a distinct lutocline, clearing of the upper water column, and increasing concentrations near the bottom (Figs. 2,5). The formation and settling of the lutocline was also confirmed by the dual frequency echo sounder (Fig. 3). Whether the condition holds for accelerating conditions and resuspension is unclear and will be investigated during the next field experiment. While a gradient Richardson number dependence appears to be appropriate for the formation of fluid mud, other formulations might be better for resuspension of fluid mud (e.g. a flux Richardson number) and require investigation.

The disaggregated inorganic grains size (DIGS) of the sediment samples shows that during the period of high freshwater discharge the suspension was unflocculated (Fig. 6). The size spectra for samples collected from the surface and bottom of the water column show that only the coarsest material settled as flood current velocities approached zero. Material below 20 μm remained in suspension indicating that little or no flocculation was occurring (Kranck, 1980). In contrast, as current speed decreased in August all particles in suspension were removed at an equal rate leading to rapid clearing of the upper water column and the creation of the fluid-mud layer. Rapid flocculation at high sediment concentration is essential for rapid deposition and the formation of fluid mud (Hill et al. 2000; Milligan et al. 2007). Observations of the DIGS over the entire sampling period show an increase in coarse material near bottom that likely leads to the development of the silty base layer found in the tidalites located along the banks of the Petitcodiac.

IMPACT/APPLICATION

Our observations will help refine our understanding of formation and maintenance of fluid mud in sheared flows and to test threshold conditions for the suppression of turbulence and carrying capacity of turbulent flows (Trowbridge and Kineke 1994). These criteria (e.g. critical Richardson number) have been used to ‘set’ a maximum concentration to drive a gravity flow. This approach has been used in the STRATAFORM and EuroSTRATAFORM projects to model the density underflow observed on the Eel Shelf (Friedrichs et al. 2000; Wright et al. 2001; Scully et al. 2002) and to predict gravity flows off the Po River (Friedrichs and Scully 2007). These models assume a critical Richardson number of $\frac{1}{4}$, but field observations of shear and density gradients in high concentrations have been limited before now.

RELATED PROJECTS

G. Kineke is a co-PI on an ONR MURI project, “Mechanisms of Fluid-Mud Interactions Under Waves.” This project includes investigators from Johns Hopkins (Dalrymple, Shen), Woods Hole Oceanographic Institution (Trowbridge, Traykovski), MIT (Liu, Mei, Yue) and Memorial University (Bentley). The major objective of this study is to examine the various mechanisms of water wave dissipation over muds and field, laboratory, and theoretical approaches will be employed.

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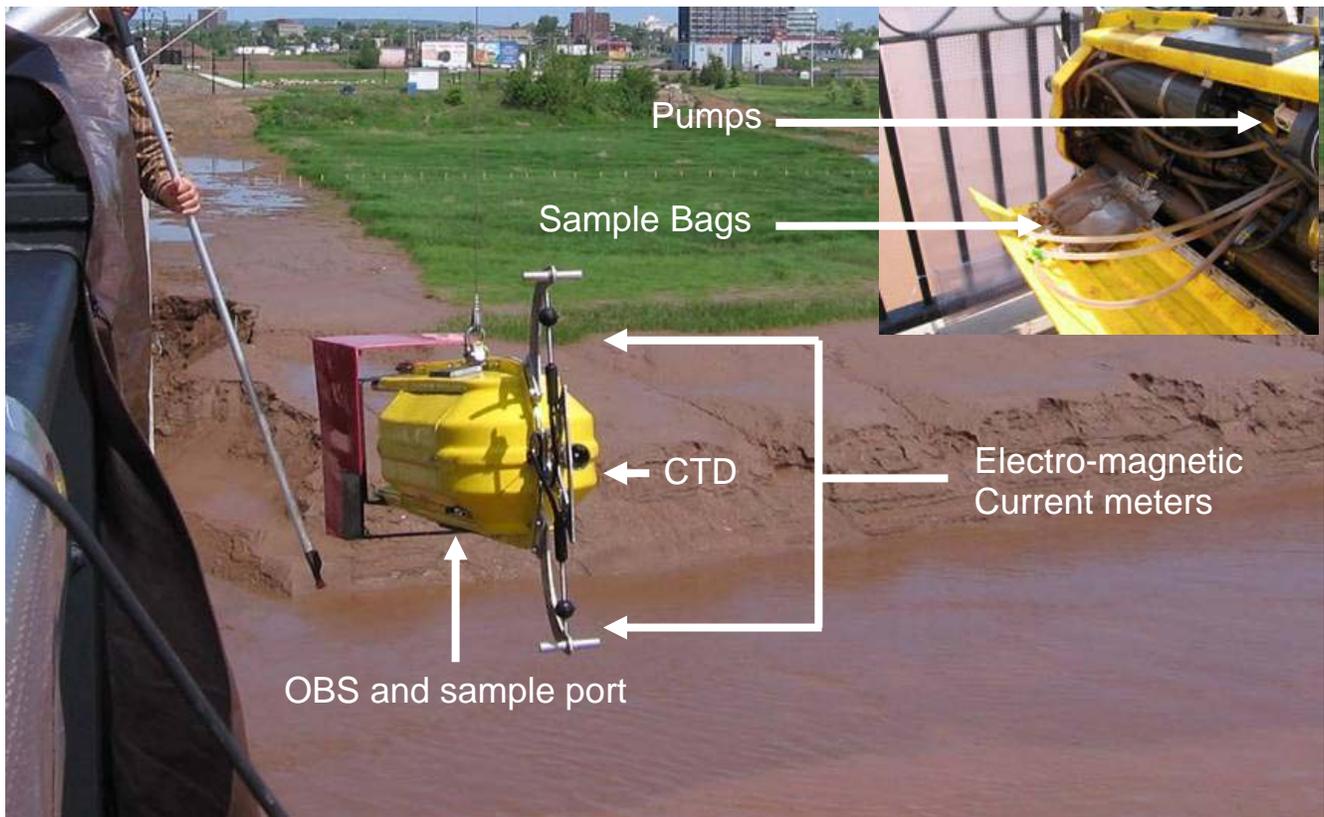


Figure 1: Image of the SUBS profiling package being deployed in the Petitcodiac River in August 2006 to observe the behaviour of fluid mud in a sheared flow. Instruments mounted on the package include a D&A optical backscatterance sensor (OBS) and sample port/pump system for collecting in situ samples, an Ocean Sensors CTD and Marsh McBirney electromagnetic current meters. Inset shows layout of pumps and sample bags inside the housing.

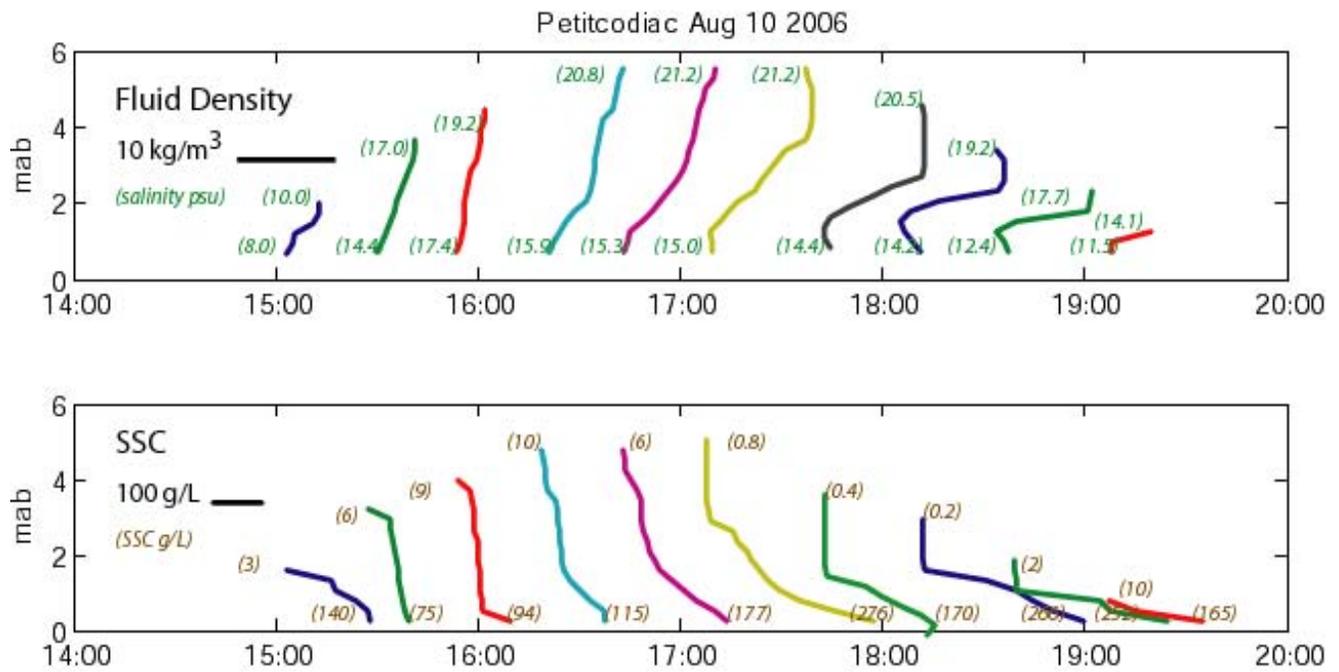


Figure 2. Water column profiles collected over a tidal cycle in the Petitcodiac River, August 10 2006. The y axis represents the distance from the bottom in meters and the x axis represents the time for each profile. The upper panel shows profiles of fluid density due to salinity and temperature. The lower panel shows profiles of suspended sediment concentration (g/L). Numbers in parentheses located at the top and bottom of each profile represent discrete values determined for those depths.

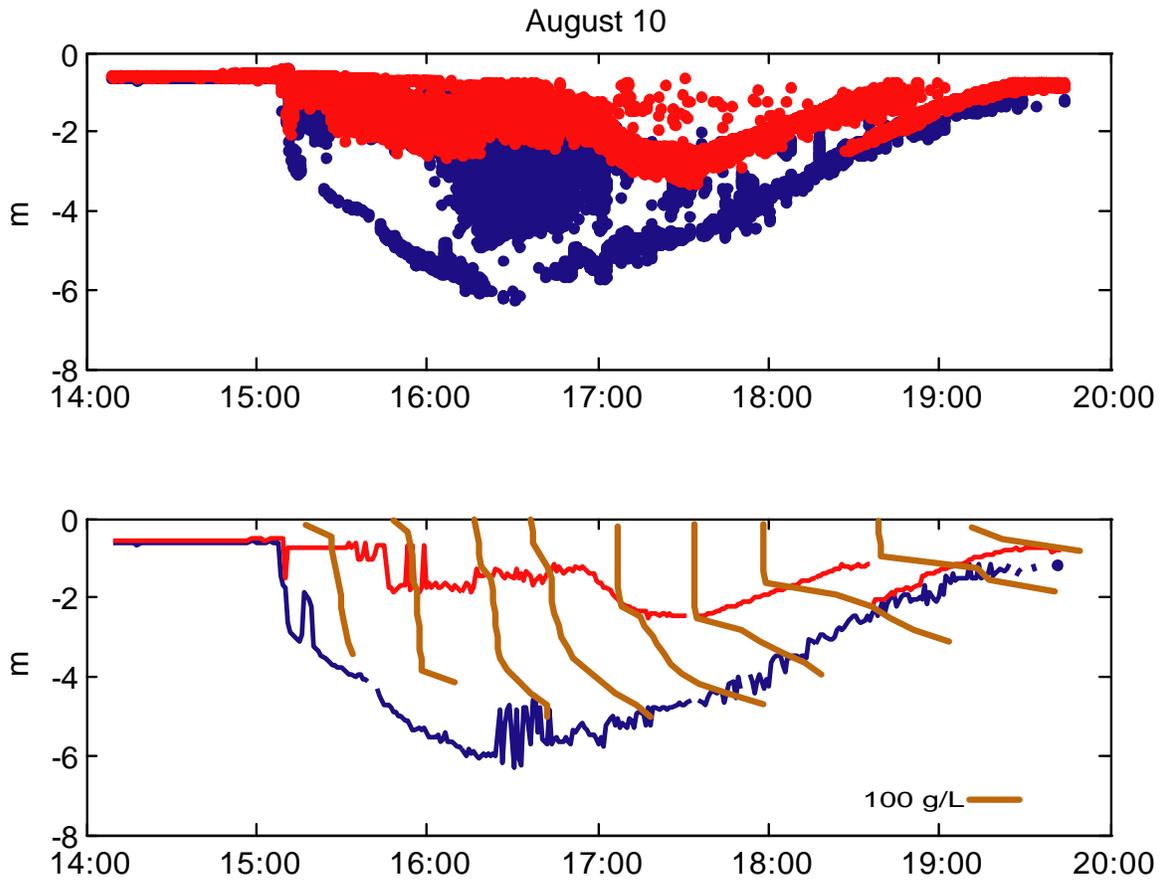


Figure 3. Dual-frequency echosounder for August 10 2006. Above: raw data with high frequency returns in red, low frequency returns in blue. Bottom: maximum depth return for 1 minute data blocks in blue, mode depth return for 1-minute data block in red. Suspended-sediment profiles from Figure 2b are superimposed showing good agreement between appearance and settling of the lutocline measured with the SUBS package and the echosounder.

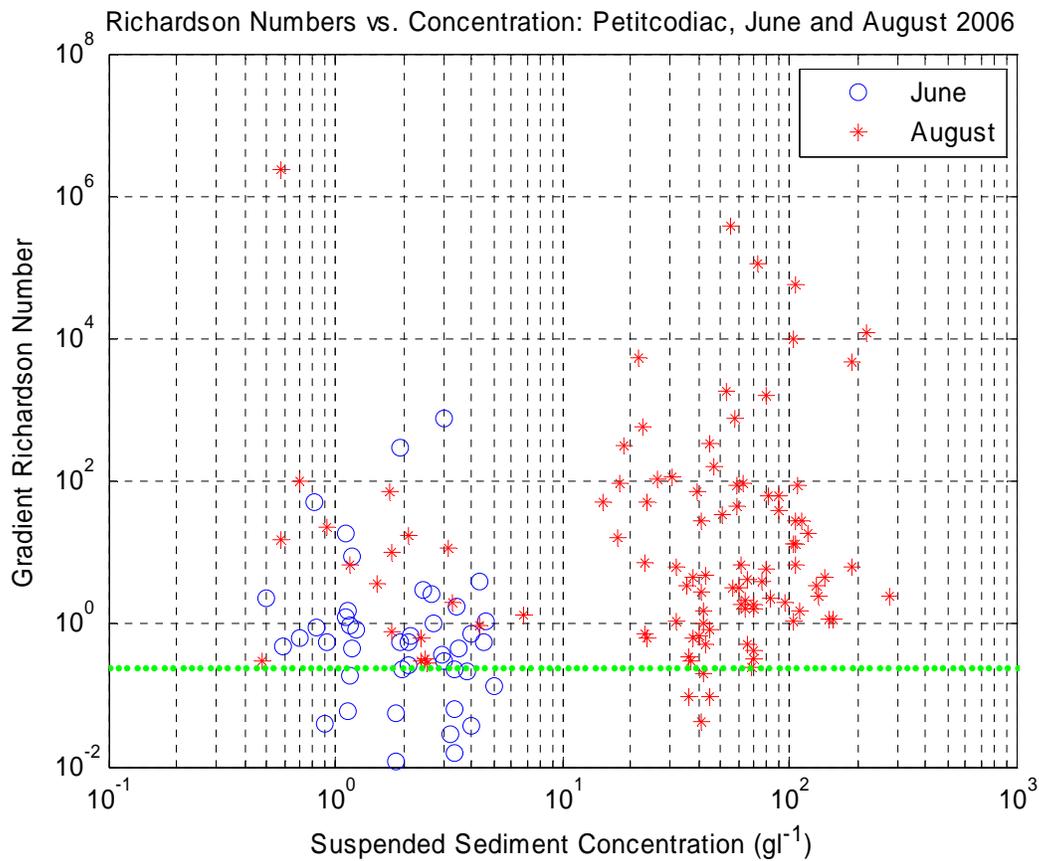


Figure 4. Graph of Gradient Richardson number vs. Suspended-Sediment Concentration for experiments conducted in June (blue open circles) and for August (red crosses) 2006. For suspended-sediment concentrations $> 10 \text{ g/L}$ Ri_g is generally greater than the critical value of $1/4$ (green dashed line). A large range of Ri_g is observed for suspended-sediment concentrations $< 1/4$, although there is a cluster around $1/4$ to 1 .

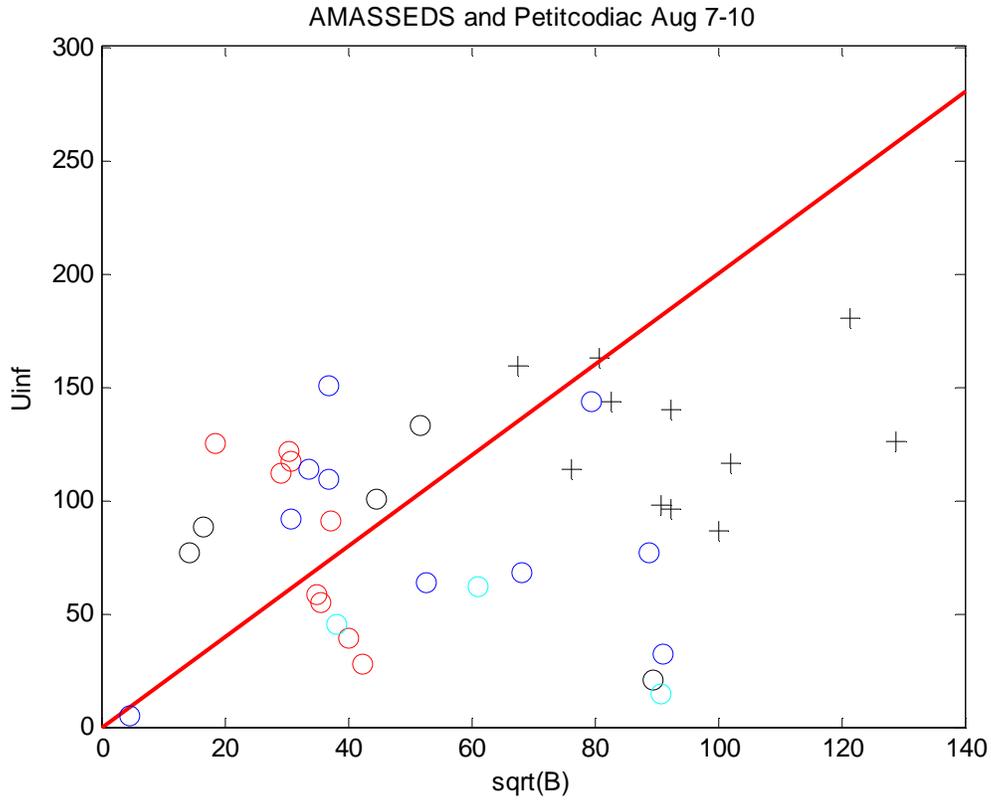


Figure 5. Free-stream velocity (U_{inf}) versus the square root of the integrated buoyancy anomaly (B). Results from FLUMES I are shown as open circles and prior results from AMASSEDS (Trowbridge and Kineke 1994) are indicated by crosses (+). Results from Aug 10 (Fig. 2) are shown in red. High-concentration bottom suspensions are expected for values below the red line.

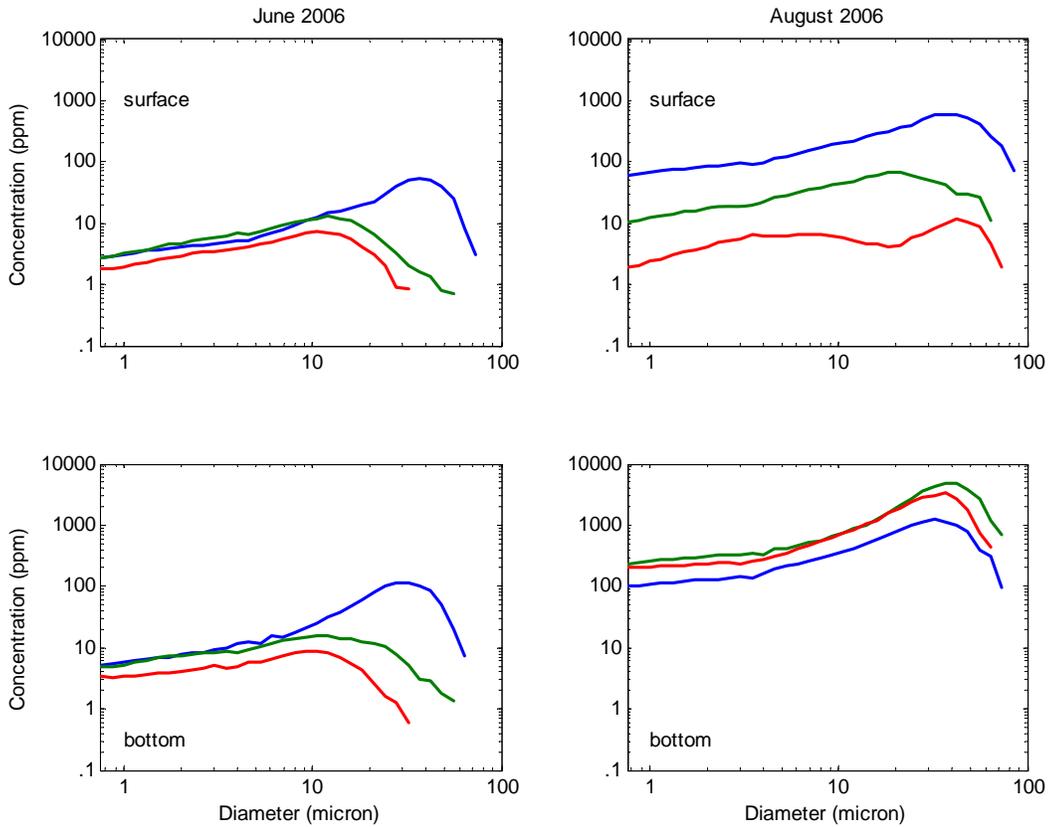


Figure 6. Disaggregated grain size for surface and bottom suspended sediment samples taken in June (fresh water) and August (salt water) for maximum (blue) and decelerating flows. The upper panels show the log of the concentration in parts per million vs log diameter for the surface samples and the lower panel show the bottom samples. The spectra for June, when no fluid mud was observed, are typical of single grain settling and those for August, when fluid mud was observed, are typical of floc settling.